



Characterizing source fingerprints and ageing processes in laboratory-generated secondary organic aerosols using proton-nuclear magnetic resonance (¹H-NMR) analysis and HPLC HULIS determination.

5 Nicola Zanca^{1,5}, Andrew T. Lambe^{2,3}, Paola Massoli², Marco Paglione¹, David R. Croasdale³, Yatish Parmar³, Emilio Tagliavini⁴, Stefania Gilardoni¹, Stefano Decesari¹

¹Institute of Atmospheric Sciences and Climate (ISAC) of the National Research Council of Italy (CNR), Bologna, 40129, Italy.

²Aerodyne Research Inc., Billerica, MA 01821, USA.

³Chemistry Department, Boston College, Chestnut Hill, MA 02467, USA.

10 ⁴Department of Chemistry "Giacomo Ciamician", University of Bologna, Bologna, 40126, Italy.

⁵Proambiente S.c.r.l., Bologna, 40129, Italy.

Corresponding author: Nicola Zanca (n.zanca@consorzioproambiente.it)

15 **Abstract.** The study of secondary organic aerosol (SOA) in laboratory settings has greatly increased our knowledge of the diverse chemical processes and environmental conditions responsible for the formation of particulate matter starting from biogenic and anthropogenic volatile compounds. However, characteristics of the different experimental setups and the way they impact the composition and the timescale of formation of SOA are still subject to debate. In this study, SOA samples were generated using a Potential Aerosol Mass (PAM) oxidation flow reactor using α -pinene, naphthalene and isoprene as precursors. The PAM reactor facilitated exploration of SOA composition over atmospherically-relevant photochemical aging time scales that are unattainable in environmental chambers. The SOA samples were analyzed using two state-of-the-art analytical techniques for SOA characterization - proton nuclear magnetic resonance (¹H-NMR) spectroscopy and HPLC determination of humic-like substances (HULIS). Results were compared with previous Aerodyne aerosol mass spectrometer (AMS) measurements. The combined ¹H-NMR, HPLC, and AMS datasets show that the composition of the studied SOA systems tend to converge to highly oxidized organic compounds upon prolonged OH exposures. Further, our ¹H-NMR findings show that only α -pinene SOA acquire spectroscopic features comparable to those of ambient OA when exposed to at least 1×10^{12} molec OH / cm³*s OH exposure, or multiple days of equivalent atmospheric OH oxidation. Over multiple days of equivalent atmospheric OH exposure, the formation of HULIS is observed in both α -pinene SOA (maximum yield = 16%) and in naphthalene SOA (maximum yield = 30%), providing evidence of the formation of humic-like polycarboxylic acids in unseeded SOA.

30 **1 Introduction**

Organic aerosol (OA) constitutes a large proportion of ambient particulate matter, affecting the Earth's radiation balance, cloud formation and human health (Hallquist et al., 2009). Understanding and simulating the concentration and composition of OA particles is one of the major challenges of modern atmospheric chemistry. In the mid 2000's, the discovery of a dominant fraction of oxidized organic compounds over primary organic compounds outside urban areas (Zhang et al., 2007; Jimenez et al., 2009), together with the understanding that the ambient organic aerosol concentrations were systematically under-predicted by existing chemical transport models (Heald et al., 2005), led to a reevaluation of the treatment of secondary organic aerosol (SOA) formation processes in chemistry and climate models. Since the model-measurement gap is mostly overcome by subjecting the particles to "oxidative aging", understanding the nature of ageing processes has become a primary objective of new generation SOA studies. Experimental findings showing the existence of highly-oxidized SOA molecular tracers with a high oxygen-to-carbon (O/C) ratio (Szmigielski et al., 2007) and molecular structures that are chemically distinct from 1st and 2nd generation oxidation products of the same precursors (Jenkin et al., 2000) provided indirect confirmation of still unknown chemical processes forming highly oxidized, low-volatility compounds.



45 The first formulations of SOA ageing into models were based on the chemistry of saturated hydrocarbon oxidation by OH, for which a step-by-step process with a slow, progressive increase of the oxidation state, along with a decrease in volatility, can be proposed (Robinson et al., 2007). The duration of such processes clearly exceeds the residence time of SOA in traditional environmental chamber experiments with equivalent atmospheric ageing times of less than 1 day. These limitations led to the emergence of oxidation flow reactors that are capable of higher integrated oxidant exposures, including the Potential Aerosol Mass (PAM) oxidation flow reactor (Kang et al., 2007; Lambe et al., 2011) and related techniques (Hall IV et al., 2013; Keller and Burtscher, 2012; Slowik et al., 2012). Recent studies suggest that flow reactor-generated SOA particles have similar composition to SOA generated in chambers (Lambe et al., 2015; Bruns et al., 2015). Modeling work further suggests that flow reactors simulate tropospheric oxidation reactions with minimal experimental artifacts (Li et al., 2015; Peng et al., 2015, 2016). Recent applications of oxidation flow reactors in field measurements showed that the maximum yields of SOA were attained at approximately 2-3 days of equivalent atmospheric OH oxidation; at higher photochemical age, SOA yields decrease substantially (Tkacik et al., 2014; Ortega et al., 2016; Palm et al., 2016). Such observations demonstrate the influence of fragmentation reactions in which oxidation leads to C-C bond cleavage with the production of highly volatile products (Kroll et al., 2009; Chacon-Madrid and Donahue 2011; Lambe et al., 2012).

60 The idea of a slow, multi-generation SOA ageing was recently challenged by recent findings from reaction chamber experiments employing modern chemical ionization mass spectrometric methods. For example, it was found that oxidized gaseous compounds with O/C > 0.7 form readily upon VOC oxidation (Ehn et al., 2012, 2014; Krechmer et al., 2015; Rissanen et al., 2014) and that even the chemical tracers of “aged” SOA can be in fact produced among 2nd generation oxidation products (Müller et al., 2012). The quantification of highly oxidized SOA compounds in reaction chambers is challenging because of significant vapor and particle wall losses (Zhang et al., 2014; Krechmer et al., 2016), but these findings suggest that SOA ageing can be much faster than previously thought (Hodzic et al., 2016). As a result of the diverse implementations of SOA schemes in models, the quantification of SOA production and concentration in the atmosphere is still highly uncertain: a recent intercomparison between 20 state-of-the-art global models showed that the estimated SOA annual production rates differ of one order of magnitude (Tsigaridis et al., 2014). These results call for more experimental observations for constraining the existing SOA parameterizations.

70 The present study focuses on laboratory production of SOA from three different precursors using a PAM reactor. The novel feature of this work is our application of two off-line analytical techniques that provide valuable insight in regards to SOA composition yet are rarely employed for SOA characterization. The first technique, ¹H-NMR spectroscopy, is a universal technique in organic chemistry. It was used to confirm the molecular structures of many SOA tracers (Finessi et al., 2014) or for following SOA reaction products in aqueous solution (Yu et al., 2011). The few examples of ¹H-NMR spectroscopy on SOA complex mixtures (Cavalli et al., 2006; Baltensperger et al., 2008; Bones et al., 2010) indicate that the technique can be very specific for distinguishing different biogenic and anthropogenic SOA systems. The present study is the first application of ¹H-NMR spectroscopy to SOA samples produced from the OH oxidation of biogenic and anthropogenic SOA. The acquisition of NMR fingerprints for fresh and aged biogenic and anthropogenic SOA can be useful also for interpreting factor analysis results obtained on a timeline of NMR spectra of ambient aerosol extracts (e.g., Paglione et al. 2014a). The second technique is a HPLC set-up for the determination of humic-like substances (HULIS). HULIS have been observed in ambient organic aerosol for nearly two decades (Havers et al., 1998; Limbeck et al., 2003), but their formation pathways aside from production in biomass burning plumes remain unclear (Graber and Rudich 2006). It is well-known that high-molecular weight oxygenated organic compounds readily form by heterogeneous reactions (Limbeck et al., 2003) or by gas-to-particle conversion (Kalberer et al., 2006), but there is little evidence for their identification with HULIS in ambient aerosols (especially if we base the definition on the chromatographic behavior, as in Baduel et al., 2009).

85 Here we focus on SOA systems generated from three distinct precursors: isoprene, α -pinene and naphthalene. Naphthalene is used as proxy for anthropogenic aromatic intermediate volatility organic compounds (IVOCs). Alpha-pinene is the most studied biogenic monoterpene due to its global importance as a biogenic SOA precursor (e.g. Pye et al., 2010), while isoprene is the most abundant biogenic VOC, accounting for 44% of global emissions (Guenter et al., 1995). The discovery of isoprene SOA is relatively recent (Claeys et al. 2004). In the presence of acidic wet aerosols, SOA originates from the heterogeneous uptake of isoprene epoxides (“IEPOX channel”, Lin et al., 2012). Since aerosol water and acidity are primarily determined by anthropogenic mineral acids, the formation of SOA from isoprene appears to be very much controlled by anthropogenic emissions. 90 On the other hand, recent experiments conducted at very low nitrogen oxide (NO) concentrations, and in the absence of seed aerosols, showed that SOA can still form from isoprene (“non-IEPOX” SOA, Krechmer et al., 2015). Such aerosols are more representative of the preindustrial world and their characterization is of paramount importance for understanding the climate radiative forcing of SOA at the global scale. Our results, obtained in the PAM reactors in absence of NO_x, are representative for non-IEPOX isoprene SOA.



2 Experimental methods

2.1 PAM oxidation flow reactor

100 The PAM oxidation flow reactor is a horizontal 13L glass cylindrical chamber that is 46 cm long x 22 cm ID. Carrier gas flows of 8.5 liter per minute (lpm) N₂ and 0.5 lpm O₂ were used, with 8.5 lpm of flow pulled through the reactor and 0.5 lpm of excess flow removed prior to the reactor. Other experimental details are fully described in Lambe et al. (2011). In this study, the PAM reactor was connected to a Scanning Mobility Particle Sizer (SMPS), an Aerodyne time-of-flight aerosol mass spectrometer (AMS), and a filter holder equipped with 47 mm (prebaked) quartz-fiber filters. SOA concentrations calculated from SMPS and/or AMS measurements averaged over filter collection times provided an estimate of the organic matter loading on the filters.

105 During the first set of experiments, α -pinene and naphthalene SOA samples were generated at multiple ageing states, by varying the UV actinic flux inside the reactor through changing the voltage applied to the lamps (Table 1). A total of five α -pinene and five naphthalene SOA samples were obtained (collection time between 3 and 20 h), with integrated OH exposures varying between 2×10^{11} and 2×10^{12} molecules/cm³*s, corresponding to a photochemical age of 1.5 to 15 days assuming a 24-hour average OH concentration of 1.5×10^6 molec cm⁻³ (Mao et al., 2009).

110 During the second set of experiments, isoprene SOA samples were generated in the reactor (Table 2). Due to the lower yields of SOA produced by isoprene oxidation, samples were collected at OH exposure of approximately 8×10^{11} molec cm⁻³ sec (corresponding to a photochemical age of 6 days) at which the maximum SOA yield is obtained (Lambe et al., 2015). The collection time was varied between 2 and 18 h.

2.2 Extraction and off-line sample characterization

115 Each filter was extracted with 5 mL of deionized ultra-pure water (Milli-Q) in a mechanical shaker for 1 h and the water extract was filtered on PTFE membranes (pore size: 0.45 μ m) in order to remove suspended particles. The water extracts were dried by rotary evaporator and were then re-dissolved in 2.15mL of D₂O: 0.65mL for proton-nuclear magnetic resonance (¹H-NMR) characterization (Decesari et al., 2000), and 1.5mL for HPLC analysis and total organic carbon (TOC) analysis (Mancinelli et al., 2007). Tests of extraction using methanol instead of water were carried out on three isoprene SOA samples. The ¹H-NMR spectra of methanol extracts were completely consistent with those obtained for the other three analyzed in deuterated water (Fig. S1), indicating that there were no specific classes of water-insoluble compounds in the isoprene SOA under the conditions used in this study. The following discussion will focus on the water-soluble fraction for which spectroscopic and chromatographic data were obtained for all three SOA systems.

2.3 NMR spectroscopy

130 The ¹H-NMR spectra were acquired at 600MHz with a Varian 600 spectrometer in a 5mm probe with 0.65mL of each sample re-dissolved in D₂O. Sodium 3-trimethylsilyl-(2,2,3,3-d₄) propionate (TSP-d₄) was used as the referred internal standard. A buffer of potassium deuterated formate/formic acid (pH~3.8) was used in the second series of experiments (isoprene SOA) to stabilize the chemical shift of hydrogen atoms in acyl functional groups, while the extracts obtained during the first experiments (α -pinene and naphthalene) were analyzed unbuffered. ¹H-NMR spectroscopy of low-concentration samples in protic solvents provides the speciation of hydrogen atoms bound to carbon atoms (H-C). On the basis of the range of frequency shifts (the chemical shift, in ppm) in which the NMR resonances occur, they can be attributed to different H-C containing functional groups (Paglione et al. 2014a).

2.4 HPLC-UV-TOC method

135 HULIS were determined using the ion exchange chromatographic method described by Mancinelli et al. (2007). A HPLC system (Agilent Model 1100) with gradient elution was used. The subsequent elution of chemical compounds bearing zero, one, two or more than two ionized groups per molecule (mainly carboxylate ions at pH 7) is monitored by an UV detector at 260 nm. Downstream of the detector, a fraction collector is programmed to sample separately “neutral compounds” (NC), “monocarboxylic acids” (MA), “dicarboxylic acids” (DA), and “polycarboxylic acids” (PA) or HULIS. The amount of WSOC recovered in each fraction is determined off-line by TOC analysis using an Analytik-Jena multi analyzer N/C (model 2100). The full HPLC-UV-TOC analytical protocol is reported in the Supplementary material. Further information on the nature of the chemical classes separated by the HPLC method based on elution tests of standard compounds, including discussion of possible misclassification, is reported by Decesari et al. (2005).

145



3.1 Results: ¹H-NMR results

3.1.1 NMR fingerprints of fresh and aged α -pinene SOA

Figure 1 shows the ¹H-NMR spectra of α -pinene SOA with increasing photochemical age. The first spectrum corresponding to a “low” SOA oxidation level is similar to reported NMR spectra of environmental chamber-generated α -pinene ozonolysis (Cavalli et al., 2006). However, the NMR fingerprint of α -pinene SOA evolves rapidly with further oxidation steps. A clear, progressive disappearing of 1st generation oxidation products (pinic and pinonic acid) with increasing O/C ratio can be observed. In the NMR spectra corresponding to a “medium” SOA oxidation level, the resonances at ~1 ppm of chemical shift, arising from one of the two gem-methyls of α -pinene, have almost disappeared. This indicates that α -pinene SOA composition evolves rapidly towards more highly-oxidized molecular structures with little resemblance to 1st generation oxidation products. At “medium” and “high” oxidation levels, the unsubstituted alkyl groups of the SOA mixture give rise to a broad Gaussian band between 1.1 and 1.8 ppm of chemical shift with a maximum at 1.4 – 1.5 ppm. The middle point position showing a slight deshielding with respect to a purely alkylic chain (~ 1.3 ppm for fatty acids) indicates the presence of electronegative groups (such as oxygen atoms) in beta or gamma position with respect to these alkyl groups. The band between 1.9 and 3 ppm, attributable to C-H groups of acyl groups (HC-C=O), also shows a transition towards structures containing more deshielded H atoms. In all cases, the most conspicuous band in this spectral region is found at 2.2 – 2.3 ppm of chemical shift, which corresponds to acetyl and acyl groups of aliphatic compounds with a low O/C ratio, like pinonic acid (R-CH₂-C=O and CH₃-(C=O)-R, where R is mainly a C_xH_{2x+1} radical). Such band persists in all SOA samples, but an additional band between 2.5 and 2.9 ppm is observed in the two samples with the highest O/C ratio, indicating that the aliphatic groups become more and more substituted by electronegative groups: the keto and carboxylic groups become spaced by no more than two methylene (or methyns) groups (x(C=O)-CH-CH-(C=O)X, where X is a generic substituent). Finally, the NMR resonances in the third important aliphatic region of alkoxy groups (CH-O), between 3.3 and 4.2 ppm of chemical shift, are always relatively small, with peak intensity at intermediate photochemical age. SOA species that contribute to NMR resonance in this region may be correlated with semivolatile, highly functionalized species that contribute to maximum SOA yields observed at intermediate OH exposures (Lambe et al., 2015).

Overall, the NMR fingerprint of α -pinene SOA is highly dependent on photochemical age, with a sharp change already at medium ageing. The most oxidized samples show spectral features that have lost any clear similarity with those of SOA sampled in reaction chambers experiments without ageing (Cavalli et al., 2006).

3.1.2 NMR fingerprints of fresh and aged naphthalene SOA

The ¹HNMR spectra of naphthalene SOA samples with increasing O/C ratio are presented in Figure 2. The extract of the least-oxidized sample (on the top) shows broad resonances between 0 and 3 ppm probably due to the effect of colloidal hydrophobic material in solution. Despite such artifact, all spectra of fresh and moderately aged SOA show clear signals from aromatic structures (region at chemical shift from 6.5 to 8.5 ppm) and alkenes (approximately between 5 and 7 ppm). The HNMR spectra of naphthalene SOA are very different than HNMR spectra of SOA produced from the OH oxidation of one-ring aromatic VOCs (Baltensperger et al., 2008), which have mostly aliphatic groups originating from ring opening reactions. Our data are in agreement with molecular speciation studies (Lee and Lane 2009), indicating that naphthalene is oxidized to form 1- or 2-ring aromatic compounds such as naphthol, and substituted benzoquinones, cinnamic acid, and phthalic acid (Chhabra et al., 2015). The chemical shift range of the main aromatic band, between 7.4 and 8.1 ppm, indicates that aromatic rings are substituted prevalently by electron-withdrawing groups, such as carbonyl and carboxyls. At moderate ageing states, a small band at 6.9 – 7.1 ppm indicates the formation of phenolic structures. The HNMR spectra of fresh naphthalene SOA show several singlets in the aromatic region, indicating a diversity of individual compounds occurring in relatively high concentrations, while moderately aged SOA show mainly the two singlets of phthalic acid. All spectra contain signals at lower chemical shifts with respect to the aromatics, mainly between 3.5 and 6.0 ppm (with the interference of the partly suppressed peak of water in the middle): several functional groups can give rise to these bands, including alkoxy-, peroxy-, esters, hemiacetals and acetals, and vinyls.

The spectrum of the most aged sample (NAPTH 1), which is also the one with the lowest SOA concentration, shows an interference from α -pinene SOA in the aliphatic region. This feature is likely due to experimental setting contamination from a previous α -pinene SOA experiment. Despite the low naphthalene SOA concentration, a broad aromatic band between 6.5 and 8.5 ppm and the same signals found between 3.5 and 6.0 ppm seen in the samples with a medium O/C ratio are still visible in this most aged naphthalene SOA spectrum. Unlike α -pinene SOA, there is no clear trend in the formation/disappearance of aromatic and aliphatic



bands with ageing, and overall the $^1\text{H-NMR}$ fingerprint of naphthalene SOA appears less sensitive to variations in the OH exposure compared to the α -pinene SOA.

3.1.3 NMR fingerprints of non-IEPOX isoprene SOA

200

The samples of isoprene SOA were obtained from the same OH exposure (corresponding approximately to a “medium” exposure in the α -pinene and naphthalene experiments) and differed only for collection time and sample quantity loaded on the filter. The isoprene SOA $^1\text{H-NMR}$ spectra profiles were all very similar (an example is provided in Figure 3). The comparison with literature data (Budisulistiorini et al. 2015) led to the unambiguous identification of 2-methyltetrols, responsible for the two singlets at 1.12 ppm (methyl H atoms of methylerythritol) and 1.13 ppm (methyl H atoms of methylthreitol) and for a series of multiplets between 3.4 and 3.9 ppm in a clear pattern. Methylerythritol is slightly more abundant (on average 55% of the sum of the two) than methylthreitol. The two methyltetrols account for 65% of the total $^1\text{H-NMR}$ signal, the rest being characterized by broad background signal with very few sharp resonances, indicating that the isoprene SOA samples are composed mainly of methyltetrols together with a significant amount of mass composed of a very complex mixture of products. The unresolved background resonances are located below the peaks of the methyltetrols, suggesting that the complex mixtures (which can include also oligomeric species) encompass molecular species (or monomers) similar to methyltetrols (at least in their C-H backbone). However, the range of chemical shifts of the background bands characterize molecular species with more electronegative groups (leading to more de-shielded H atoms) than methyltetrols: the band of methyl protons extends to 1.7 ppm (respect to 1.12-1.13 of methyltetrols) and the band of alcoxyl groups (HC-O) extends to 4.3 ppm. The results of Liu et al. (2016), indicating that non-IEPOX isoprene SOA include peroxide-equivalents of methyltetrols, are in agreement with these findings. However, Liu et al. (2016) do not report the presence of carboxylic or keto groups, while our data clearly indicate that these (and/or other acyl groups) are found in the unresolved mixtures of non-IEPOX isoprene SOA, and are responsible for the signal band between 2.0 and 2.6 ppm. Still, this band is much less intense than that of alcoxyls, which is opposite to what observed for α -pinene SOA where acyls are by far the main oxygenated aliphatic functional group. Thus, $^1\text{H-NMR}$ spectroscopy provides distinct fingerprint for isoprene and monoterpene SOA.

205

210

215

220

3.2 HPLC results

The HPLC analysis of fresh α -pinene SOA extracts show the presence of compounds unretained by ion-exchange columns (neutral compounds) or weakly retained (mono- and di-acids) with a small contribution from compounds having a high retention factor (polyacids, PA, or HULIS), in agreement with previous results obtained from α -pinene SOA samples generated in environmental chamber experiments (unpublished data). The HULIS content increases only moderately with ageing, while the yield/fraction of di-acids increases significantly with respect to mono-acids and neutral/basic compounds (Figure 4). With increasing photochemical age, the total organic carbon (TOC) mass fractions of mono-acids, di-acids and HULIS increase from 20% to 34% and 7% to 16%, respectively, whereas the mass fraction of neutral compounds decreases from 19% to 9%. These results are in qualitative agreement with the known chemistry of α -pinene SOA, in which mono- and dicarboxylic acids are the most characteristic condensable 1st generation products (Jenkin et al., 2000; Jaoui and Kamens 2001) while tricarboxylic acids such as 3-methyl-1,2,3-butanetricarboxylic acid or pinyl-diterpenyl ester (Szmigielski et al., 2007; Yasmeeen et al., 2010) are present in lesser amounts and can contribute the observed concentrations of HULIS in this study.

230

The HPLC fractionation of naphthalene SOA (Figure 5) shows that fresh samples are characterized by a mixture of neutral compounds and mono- and di-acids, completely consistent with the molecular compositions reported in the literature (Lee and Lane 2009). However, a net increase in acidic compounds with photochemical age can be clearly observed. The HULIS content, initially small, increases substantially and progressively with ageing. With increasing photochemical age, the TOC mass fraction of mono- and di-acids decreases from 33% to 18% and the fraction of PA/HULIS increases from 11% to 30%. This is the first evidence of HULIS formation (determined with the ion-exchange method) in laboratory-generated SOA. The formation of polyacidic molecules with three or more carboxylic groups implies the opening of the second aromatic ring in the naphthalene precursor backbone, and/or oligomerization reactions. In both cases, products of such oxidation reactions cannot be explained by current naphthalene SOA molecular speciation studies (e.g. Kautzman et al., 2010).

235

240

Finally, the HPLC analysis of isoprene SOA shows that neutral compounds (NC) were dominant in all sample extracts (Figure 6). NC accounted for 59% of the TOC content of the sum of the HPLC fractions. The second most abundant fraction (34%) are the mono-acids, while diacids accounted for a much smaller fraction of TOC, and to polyacids were almost absent (ca. 1% of the

245



TOC). The dominance of NC is consistent with the high yield methyltetrols and their analogues (see Section 3.1.3). Assuming that the distribution of NMR functional groups approximately reflects their carbon content, methyltetrols (accounting for 65% of the total NMR signal) can account for the whole of the HPLC neutral compounds, and, as a corollary, the complex mixtures of products detected by unresolved bands by NMR spectroscopy correspond to the mono- and di-acids in HPLC. As already noticed in the previous section, the NMR analysis shows indeed the occurrence of acyl groups which indicate/support the presence of carboxylic acids. We cannot exclude, however, possible misclassification of some neutral compounds into the monoacid fraction, as already observed for some dicarbonyls (Decesari et al. 2005).

4 Discussion and conclusions

In this section, the NMR and HPLC results obtained for the isoprene, α -pinene and naphthalene SOA systems are compared with ambient OA samples. First, we investigated the similarity between the $^1\text{H-NMR}$ spectral profiles of SOA with those “typical” of ambient non-biomass-burning WSOC. For this purpose we used one sample of PM_{10} from the rural Po Valley in summertime, collected during the 2012 PEGASOS field campaign (Sandrini et al., 2016). The PEGASOS and laboratory SOA spectra were binned to 400 points in order to remove the variability in chemical shifts due to, e.g., different pH conditions during the analyses of the samples. Figure 7 shows the correlation between the SOA spectra and the reference spectrum of ambient WSOC. There is a good correlation ($0.62 < r < 0.92$) between the NMR spectra of α -pinene SOA, at all oxidation levels, with the spectrum of the PM_{10} sample, whereas a low correlation is found for the isoprene SOA ($0.26 < r < 0.30$); the naphthalene SOA spectra exhibit zero or negative correlations ($-0.15 < r < -0.02$). This result is somewhat expected if we consider that ambient water-soluble aerosols are characterized by acyl functional groups (HC-C=O) in higher concentrations with respect to alkoxy groups (HC-O) and by a smaller aromatic content (Decesari et al., 2007), with a functional group pattern that is well reproduced by α -pinene SOA and not by non-IEPOX SOA and by naphthalene SOA. Clearly, naphthalene SOA *alone*, with hydrogen-to-carbon (H/C) ratio less than 0.9 due to relatively high aromatic content (Lambe et al. 2011, 2015), does not mimic ambient OA *bulk* composition. It should be noted, however, that naphthalene and other polyaromatic hydrocarbons are co-emitted with many other anthropogenic IVOCs and VOCs including aliphatic compounds in the real atmosphere; therefore the contribution of naphthalene SOA could be simply masked by the contribution of aliphatic IVOC SOA in the $^1\text{H-NMR}$ spectra of ambient WSOC. When limiting the correlation analysis of Figure 7 to the aromatic and vinyl region of the spectra (> 6 ppm), Pearson r coefficients of 0.49 to 0.58 are found between the naphthalene SOA and the ambient WSOC spectra, while small values (between -0.2 and 0.4) are found for the α -pinene and isoprene SOA. This result suggests that SOA produced from naphthalene or similar precursors, including many other ring-retaining oxidation products (Section 3.1.2) can explain the presence of aromatic moieties in ambient water-soluble aerosols in areas not affected by biomass burning emissions.

When considering the full spectral range, the H-NMR spectra of α -pinene SOA most closely mimic the functional group distributions of the ambient WSOC sample obtained in PEGASOS. Interestingly, similarity between H-NMR spectra of α -pinene SOA and Po Valley WSOC increases with increasing photochemical age. For the most oxidized α -pinene SOA samples, the functional group composition, characterized by polysubstituted aliphatic compounds rich of acyls (carboxylic or keto groups), overlaps well with that of ambient WSOC. A good fit between α -pinene SOA and ambient WSOC mass spectral features is already achieved at medium oxidation conditions, in agreement with the results of Lambe et al. (2011) showing that the correlation between the AMS spectra of PAM-generated α -pinene SOA with the spectra of ambient OOA increases up to an exposure of 1×10^{12} OH molec cm^{-3}s and remains rather stable afterwards. These results suggest that NMR-traced ageing processes reflect the same chemical mechanisms already studied using high-resolution AMS techniques. However, the correlation coefficients shown in Figure 7 for the NMR spectra of α -pinene vs. ambient WSOC ($r^2 = 0.39$ to 0.84) are smaller than those between the HR-ToF-AMS spectra of PAM-generated SOA vs. ambient OOA ($r^2 = 0.7$ to 0.9) (Fig. 9 in Lambe et al., 2011). Apparently, the AMS features of ambient OA are more easily reproduced by PAM experiments than the NMR composition, or, in other words, NMR spectroscopy exhibits a greater selectivity for the OA components than AMS. Specifically, $^1\text{H-NMR}$ spectroscopy was able to resolve significant changes in composition of α -pinene SOA with photochemical ageing in great detail, especially at an OH exposure of $\sim 1.1 \times 10^{12}$ molec OH / cm^3s equivalent to multiple days of atmospheric ageing.

A comparison of fresh and aged SOA with ambient WSOC samples with respect to the HPLC fraction distributions is reported in Figure 8. The distribution of neutral vs. acidic classes of compounds in ambient WSOC refers to average of the samples collected at a rural background station at Cabauw in the Netherlands (Paglione et al. 2014b). The HULIS contribution in these samples varied between 15 and 20%, in line with previous results obtained in the Po Valley (Mancinelli et al., 2007), but lower than in biomass burning aerosol samples (Decesari et al., 2006). The α -pinene SOA generated in the PAM reactor at high photochemical age are characterized by a HULIS amount similar to that of Cabauw samples, while the polyacidic content of aged naphthalene SOA is higher than in the ambient samples. These results demonstrate that laboratory experiments of SOA formation can generate



complex mixtures of products with the same chromatographic properties of HULIS provided a sufficient extent of photochemical aging using the PAM reactor or related techniques. The HULIS fraction of WSOC is in fact proportional to the AMS f44 for SOA (integrated over the filter sampling times) (Figure 9) irrespectively of precursor type. Therefore, the formation of polycarboxylic acids determined by the HPLC technique follows the same positive trend in concentrations of the AMS proxy for C(O)OH groups with increasing OH exposure. This is in contrast with the numerous observations of rapid formation of SOA oligomers during reaction chamber experiments (Kalberer et al., 2006; Reemtsma et al., 2006), indicating that oligomers do not account for chromatographically-defined HULIS. A survey of the laboratory studies on the formation of humic material in secondary aerosol shows that evidence for the formation of polycarboxylic acids comes from the reaction of phenolic compounds in the presence of particulate water (Hoffer et al., 2004), while little is known for unseeded, dry gas-to-particle formation experiments. With the exception of the very preliminary data reported by Baltensperger et al. (2008), our results – to our best knowledge – are the first showing the occurrence of HULIS *sensu strictu* in monoterpene and aromatic hydrocarbon SOA, and these HULIS are clearly shown to be a product of photochemical ageing.

In conclusion, we observed that OA ageing reactions in the PAM reactor produces water-soluble compounds of high complexity but with spectroscopic and chromatographic properties that converge towards those characteristic of ambient OA. Specifically, a good correlation between ambient HPLC/HNMR samples and aged α -pinene SOA was observed in respect to HULIS content and NMR functional group distribution, while aged aromatic IVOC SOA show clear potential for HULIS formation. The isoprene SOA samples do not show compositional features with a clear overlap with those of ambient WSOC obtained in the Cabauw and Po Valley samples that are representative of continental polluted atmospheres, but should serve as useful reference spectra for future studies/environments impacted by non-IEPOX isoprene SOA.

Acknowledgements

A.T. Lambe and P. Massoli acknowledge support by the Atmospheric Chemistry Program of the U.S. National Science Foundation under grants AGS-1536939, AGS-1537446 and by the U.S. Office of Science (BER), Department of Energy (Atmospheric Systems Research) under grants DE-SC0006980 and DE-SC0011935. We thank Manjula Canagaratna (ARI), Douglas Worsnop (ARI), Timothy Onasch (BC/ARI) and Paul Davidovits (BC) for helpful discussions. S. Decesari, S. Gilardoni, M. Paglione and N. Zanca acknowledge funding from the European FP7 project BACCHUS (grant agreement No. 49 990 603445). Dr. Fabio Moretti, formerly at the Department of Chemistry of the University of Bologna, Prof. Andrea Mazzanti and Dr. Alessandra Petrolini of the Department of Industrial Chemistry of the University of Bologna are also greatly acknowledged for support with the NMR analyses.

References

- Baduel, C., Voisin, D., and Jaffrezo, J.-L.: Comparison of analytical methods for Humic Like Substances (HULIS) measurements in atmospheric particles, *Atmos. Chem. Phys.*, 9, 5949–5962, 2009.
- Baltensperger, U., Dommen, J., Alfarra, M. R., Duplissy, J., Gaeggeler, K., Metzger, A., Facchini, M. C., Decesari, S., Finessi, E., Reinnig, C., Schott, M., Warnke, J., Hoffmann, T., Klatzer, B., Puxbaum, H., Geiser, M., Savi, M., Lang, D., Kalberer, M., Geiser, T.: Combined determination of the chemical composition and of health effects of secondary organic aerosols: the POLYSOA project, *J. Aerosol Med Pulm Drug Deliv.*, 21, 145–154, 2008.
- Bones, D. L., Henricksen, D. K., Mang, S. A., Gonsior, M., Bateman, A. P., Nguyen, T. B., Cooper, W. J., and Nizkorodov, S. A.: Appearance of strong absorbers and fluorophores in limonene-O₃ secondary organic aerosol due to NH₄⁺ - mediated chemical aging over long time scales, *J. Geophys. Res.*, 115, doi:10.1029/2009JD012864, 2010.
- Bruns, E. A., El Haddad, I., Keller, A., Klein, F., Kumar, N. K., Pieber, S. M., Corbin, J. C., Slowik, J. G., Brune, W. H., Baltensperger, U., and Prévôt, A. S. H.: Inter-comparison of laboratory smog chamber and flow reactor systems on organic aerosol yield and composition, *Atmos. Meas. Tech.*, 8, 2315–2332, 2015.
- Budisulistiorini, S. H., Li, X., Bairai, S. T., Renfro, J., Liu, Y., Liu, Y. J., Mckinney, K. A., Martin, S. T., McNeill, V. F., Pye, H. O. T., Neff, M. E., Stone, E. A., Mueller, S., Knote, C., Shaw, S. L., Zhang, Z., Gold, A., Surratt, J. D.: Examining the Effects of



- 350 Anthropogenic Emissions on Isoprene-Derived Secondary Organic Aerosol Formation During the 2013 Southern Oxidant and Aerosol Study (SOAS) at the Look Rock, Tennessee, Ground Site. *Atmos. Chem. & Phys.*, 15, 8871–8888, 2015.
- Cavalli, F., Facchini, M. C., Decesari, S., Emblico, L., Mircea, M., Jensen, N. R., and Fuzzi, S.: Size-segregated aerosol chemical composition at a boreal site in southern Finland, during the QUEST project, *Atmos. Chem. Phys.*, 6, 993–1002, doi:10.5194/acp-6-993-2006, 2006.
- 355 Chacon-Madrid, H. J., and Donahue, N. M.: Fragmentation vs. functionalization: chemical aging and organic aerosol formation, *Atmos. Chem. Phys.*, 11, 10553–10563, 2011.
- Chhabra, P. S., Lambe, A. T., Canagaratna, M. R., Stark, H. J., Jayne, J. T., Onasch, T. B., Davidovits, P., Kimmel, J. R. and Worsnop, D. R.: Application of high-resolution time-of-flight chemical ionization mass spectrometry measurements to estimate volatility distributions of α -pinene and naphthalene oxidation products, *Atmos. Meas. Tech.*, 8, 1–18, 2015.
- 360 Claeys, M., Graham, B., Vas, G., Wang, W., Vermeylen, R., Pashynska, V., Cafmeyer, J., Guyon, P., Andreae, M. O., Artaxo, P., and Maenhaut, W.: Formation of secondary organic aerosols through photooxidation of isoprene, *Science*, 303, 1173, 2004.
- 365 Decesari, S., Facchini, M. C., Fuzzi, S., and Tagliavini, E.: Characterization of water-soluble organic compounds in atmospheric aerosol: A new approach, *J. Geophys. Res.*, 105, 1481–1489, 2000.
- Decesari, S., Moretti, F., Fuzzi, S., Facchini, M. C., Tagliavini, E.: Comment on “On the use of anion exchange chromatography for the characterization of water soluble organic carbon” by H. Chang et al., *Geophys. Res. Lett.*, 32, doi:10.1029/2005GL023826, 2005.
- 370 Decesari, S., Fuzzi, S., Facchini, M. C., Mircea, M., Emblico, L., Cavalli, F., Maenhaut, W., Chi, X., Schkolnik, G., Falkovich, A., Rudich, Y., Claeys, M., Pashynska, V., Vas, G., Kourtchev, I., Vermeylen, R., Hoffer, A., Andreae, M. O., Tagliavini, E., Moretti, F., and Artaxo P.: Characterization of the organic composition of aerosols from Rondônia, Brazil, during the LBA-SMOCC 2002 experiment and its representation through model compounds, *Atmos. Chem. Phys.*, 6, 375–402, 2006.
- Decesari, S., Mircea, M., Cavalli, F., Fuzzi, S., Moretti, F., Tagliavini, E., and Facchini, M. C.: Source Attribution of Water-Soluble Organic Aerosol by Nuclear Magnetic Resonance Spectroscopy, *Environ. Sci. Technol.*, 41, 2479–2484, 2007.
- 380 Ehn, M., Kleist, E., Junninen, H., Petäjä, T., Lönn, G., Schobesberger, S., Dal Maso, M., Trimborn, A., Kulmala, M., Worsnop, D. R., Wahner, A., Wildt, J., and Mentel, T. F.: Gas phase formation of extremely oxidized pinene reaction products in chamber and ambient air, *Atmos. Chem. Phys.* 12, 5113–5127, 2012.
- Finessi, E., Lidster, R. T., Whiting, F., Elliott, T., Alfarrá, M. R., McFiggans, G. B., and Hamilton, J. F.: Improving the quantification of secondary organic aerosol using a microflow reactor coupled to HPLC-MS and NMR to manufacture ad hoc calibration standards, *Anal. Chem.*, 86, 11238–11245, 2014.
- 385 Guenther, A., Hewitt, C. N., Erickson, D., Fall, R., Geron, C., Graedel, T., Harley, P., Klinger, L., Lerdau, M., McKay, W. A., Pierce, T., Scholes, B., Steinbrecher, R., Tallamraju, R., Taylor, J., Zimmerman, P.: A global model of natural volatile organic compound emissions, *J. Geophys. Res.*, 100, 8873–8892, doi:10.1029/94JD02950, 1995.
- 390 Graber, E. R., and Rudich, Y.: Atmospheric HULIS: How humic-like are they? A comprehensive and critical review, *Atmos. Chem. Phys.*, 6, 729–753, 2006.
- 395 Hall IV, W. A., Pennington, M. R., and Johnston, M. V.: Molecular transformations accompanying the aging of laboratory secondary organic aerosol, *Environ. Sci. Technol.*, 47, 2230–2237, 2013.
- Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue, N. M., George, C., and Goldstein, A. H. et al.: The formation, properties and impact of secondary organic aerosol: current and emerging issues, *Atmos. Chem. Phys.*, 9, 5155–5236, 2009.
- 400 Havers, N., Burba, P., Lambertm, J. and Klockow, D.: Spectroscopic characterization of humic-like substances in airborne particulate matter, *J. Atmos. Chem.*, 29, 45–54, 1998.
- 405 Heald, C. L., Jacob, D. J., Park, R. J., Russell, L. M., Huebert, B. J., Seinfeld, J. H., Liao, H., and Weber, R. J.: A large organic aerosol source in the free troposphere missing from current models, *Geophys. Res. Lett.*, 32, L18809, doi:10.1029/2005GL023831, 2005.



- 410 Hodzic, A., Kasibhatla, P. S., Jo, D. S., Cappa, C. D., Jimenez, J. L., Madronich, S., and Park, R. J.: Rethinking the global secondary organic aerosol (SOA) budget: stronger production, faster removal, shorter lifetime, *Atmos. Chem. Phys.*, 16, 7917–7941, 2016.
- Hoffer, A., Kiss, G., Blazso, M., and Gelencsér, A.: Chemical characterization of humic-like substances (HULIS) formed from a lignin-type precursor in model cloud water, *Geophys. Res. Lett.*, 31, L06115, 2004.
- 415 Jaoui, M., and Kamens, R. M.: Mass balance of gaseous and particulate products analysis from α -pinene/ NO_x /air in the presence of natural sunlight, *J. Geophys. Res.*, 106, 12,541–12,558, 2001.
- 420 Jenkin, M. E., Shallcross, D. E., Harvey, J. N.: Development and application of a possible mechanism for the generation of cis-pinonic acid from the ozonolysis of α - and β -pinene, *Atmos. Environ.*, 34, 2837–2850, 2000.
- Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., et al.: Evolution of Organic Aerosols in the Atmosphere, *Science*, 326, 1525–1528, doi: 10.1126/science.1180353, 2009.
- 425 Kalberer, M., Sax, M., Samburova, V.: Molecular size evolution of oligomers in organic aerosols collected in urban atmospheres and generated in a smog chamber, *Environ. Sci. Technol.*, 40, 5917–5922, 2006.
- Kang, E., Root, M. J., Toohey, D. W., et al.: Introducing the concept of Potential Aerosol Mass (PAM), *Atmos. Chem. Phys.*, 7, 5727–5744, 2007.
- 430 Kautzman, K. E., Surratt, J. D., Chan, M. N., Chan, A. W. H., Hersey, S. P., Chhabra, P. S., Dalleska, N. F., Wennberg, P. O., Flagan R. C., and Seinfeld, J. H.: Chemical composition of gas- and aerosol-phase products from the photooxidation of naphthalene, *J. of Phys. Chem. A*, 114, 913–934, 2010.
- 435 Keller, A. and Burtscher, H.: A continuous photo-oxidation flow reactor for a defined measurement of the SOA formation potential of wood burning emissions, *J. Aerosol Sci.*, 49, 9–20, 2012.
- 440 Krechmer, J. E., Coggon, M. M., Massoli, P., Nguyen, T. B., Crouse, J. D., Hu, W., Day, D. A., Tyndall, G. S., Henze, D. K., Rivera-Rios, J. C., Nowak, J. B., Kimmel, J. R., Mauldin, R. L., Stark, H., Jayne, J. T., Sipilä, M., Junninen, H., Clair, J. M. St., Zhang, X., Feiner, P. A., Zhang, L., Miller, D. O., Brune, W. H., Keutsch, F. N., Wennberg, P. O., Seinfeld, J. H., Worsnop, D. R., Jimenez, J. L., and Canagaratna, M. R.: Formation of Low Volatility Organic Compounds and Secondary Organic Aerosol from Isoprene Hydroxyhydroperoxide Low-NO Oxidation, *Environ. Sci. Technol.* 49, 10330–10339, 2015.
- 445 Krechmer, J. E., Pagonis, D., Ziemann, P. J., and Jimenez, J. L.: Quantification of Gas-Wall Partitioning in Teflon Environmental Chambers Using Rapid Bursts of Low-Volatility Oxidized Species Generated in Situ, *Environ. Sci. Technol.*, 50 (11), 5757–5765, doi: 10.1021/acs.est.6b00606, 2016.
- 450 Kroll, J. H., Smith, J. D., Che, D. L., Kessler, S. H., Worsnop, D. R., and Wilson, K. R.: Measurement of fragmentation and functionalization pathways in the heterogeneous oxidation of oxidized organic aerosol. *Phys. Chem. Chem. Phys.*, 11, 8005–8014, 2009.
- 455 Lambe A. T., Ahern A. T., Williams L. R., Slowik J. G., Wong J. P. S., Abbatt J. P. D., Brune W. H., Ng N. L., Wright J. P., Croasdale D. R., Worsnop D. R., Davidovits P., and Onasch T. B.: Characterization of aerosol photooxidation flow reactors: heterogeneous oxidation, secondary organic aerosol formation and cloud condensation nuclei activity measurements, *Atmos. Meas. Tech.*, 4, 445–461, 2011.
- 460 Lambe, A. T., Onasch, T. B., Croasdale, D. R., Wright, J. P., Martin, A. T., Franklin, J. P., Massoli, P., Kroll, J. H., Canagaratna, M. R., Brune, W. H., Worsnop, D. R., and Davidovits, P.: Transitions from functionalization to fragmentation reactions of secondary organic aerosol (SOA) generated from the laboratory OH oxidation of alkane precursors, *Environ. Sci. and Technol.*, 46, 5430–5437, 2012.
- 465 Lee, J. Y., and Lane, D. A.: Unique products from the reaction of naphthalene with the hydroxyl radical, *Atmos. Environ.*, 43, 4886–4893, 2009.



- 470 Li, R., Palm, B. B., Ortega, A. M., Hlywiak, J., Hu, W., Peng, Z., Day, D. A., Knote, C., Brune, W. H., de Gouw, J. A., and Jimenez, J. L.: Modeling the Radical Chemistry in an Oxidation Flow Reactor: Radical Formation and Recycling, Sensitivities, and the OH Exposure Estimation Equation, *J. Phys. Chem. A*, 119, 4418–4432, 2015.
- Limbeck, A., Kulmala, M., Puxbaum, H.: Secondary organic aerosol formation in the atmosphere via heterogeneous reaction of gaseous isoprene on acidic particles, *Geophys. Res. Lett.*, 30, 1996, doi:10.1029/2003GL017738, 2003.
- 475 Lin, Y.-H., Zhang, Z., Docherty, K. S., Zhang, H., Budisulistiorini, S. H., Rubitschun, C. L., Shaw, S. L., Knipping, E. M., Edgerton, E. S., Kleindienst, T. E., Gold, A., Surratt, J. D.: Isoprene Epoxydiols as Precursors to Secondary Organic Aerosol Formation: Acid-Catalyzed Reactive Uptake Studies with Authentic Compounds, *Environ. Sci. Technol.* 46 (1), 250-258, 2012.
- 480 Liu, J., D'Ambro, E. L., Lee, B. H., Lopez-Hilfiker, F. D., Zaveri, R. A., Rivera-Rios, J. C., Keutsch, F. N., Iyer, S., Kurten, T., Zhang, Z., Gold, A., Surratt, J. D., Shilling, J. E., and Thornton, J. A.: Efficient Isoprene Secondary Organic Aerosol Formation from a Non-IEPOX Pathway, *Environ. Sci. Technol.*, 50, 9872–9880, 2016.
- Mancinelli, V., Rinaldi, M., Finessi, E., Emblico, L., Mircea, M., Fuzzi, S., Facchini, M. C., Decesari, S.: An anion-exchange high-performance liquid chromatography method coupled to total organic carbon determination for the analysis of water-soluble organic aerosols, *J. of Chrom.*, 1149, 385-389, 2007.
- 485 Mao, J., Ren, X., Brune, W. H., Olson, J. R., Crawford, J. H., Fried, A., Huey, L. G., Cohen, R. C., Heikes, B., Singh, H. B., Blake, D. R., Sachse, G. W., Diskin, G. S., Hall, S. R., and Shetter, R. E.: Airborne measurement of OH reactivity during INTEX-B, *Atmos. Chem. Phys.*, 9, 163–173, 2009.
- 490 Müller, L., Reinnig, M.-C., Naumann, K. H., Saathoff, H., Mentel, T. F., Donahue, N. M., and Hoffmann, T.: Formation of 3-methyl-1,2,3-butanetricarboxylic acid via gas phase oxidation of pinonic acid – a mass spectrometric study of SOA aging, *Atmos. Chem. Phys.*, 12, 1483–1496, 2012.
- 495 Ortega, A. M., Hayes, P. L., Peng, Z., Palm, B. B., Hu, W., Day, D. A., Li, R., Cubison, M. J., Brune, W. H., Graus, M., Warneke, C., Gilman, J. B., Kuster, W. C., Gouw, J. de, Gutiérrez-Montes, C., and Jimenez, J. L.: Real-time measurements of secondary organic aerosol formation and aging from ambient air in an oxidation flow reactor in the Los Angeles area, *Atmos. Chem. Phys.*, 16, 7411-7433, 2016.
- 500 Paglione, M., Saarikoski, S., Carbone, S., Hillamo, R., Facchini, M. C., Finessi, E., Giulianelli, L., Carbone, C., Fuzzi, S., Moretti, F., Tagliavini, E., Swietlicki, E., Stenström, K. E., Prévôt, A. S. H., Massoli, P., Canaragatna, M., Worsnop, D. and Decesari, S.: Primary and secondary biomass burning aerosols determined by proton nuclear magnetic resonance (¹H-NMR) spectroscopy during the 2008 EUCAARI campaign in the Po Valley (Italy), *Atmos. Chem. Phys.*, 14, 5089-5110, 2014a.
- 505 Paglione, M., Kiendler-Scharr, A., Mensah, A. A., Finessi, E., Giulianelli, L., Sandrini, S., Facchini, M. C., Fuzzi, S., Schlag, P., Piazzalunga, A., Tagliavini, E., Henzing, J. S., and Decesari, S.: Identification of humic-like substances (HULIS) in oxygenated organic aerosols using NMR and AMS factor analyses and liquid chromatographic techniques, *Atmos. Chem. Phys.*, 14, 25-45, doi:10.5194/acp-14-25-2014, 2014b.
- 510 Palm, B. B., Campuzano-Jost, P., Ortega, A. M., Day, D. A., Kaser, L., Jud, W., Karl, T., Hansel, A., Hunter, J. F., Cross, E. S., Kroll, J. H., Peng, Z., Brune, W. H., and Jimenez, J. L.: In situ secondary organic aerosol formation from ambient pine forest air using an oxidation flow reactor, *Atmos. Chem. Phys.*, 16, 2943–2970, 2016.
- 515 Peng, Z., Day, D. A., Stark, H., Li, R., Lee-Taylor, J., Palm, B. B., Brune, W. H., and Jimenez, J. L.: HOx radical chemistry in oxidation flow reactors with low-pressure mercury lamps systematically examined by modeling, *Atmos. Meas. Tech.*, 8, 4863–4890, 2015.
- Peng, Z., Day, D. A., Ortega, A. M., Palm, B. B., Hu, W., Stark, H., Li, R., Tsigaridis, K., Brune, W. H., and Jimenez, J. L.: Non-OH chemistry in oxidation flow reactors for the study of atmospheric chemistry systematically examined by modeling, *Atmos. Chem. Phys.*, 16, 4283–4305, 2016.
- 520 Pye, H. O. T., Chan, A. W. H., Barkley, M. P., and Seinfeld, J. H.: Global modeling of organic aerosol: the importance of reactive nitrogen (NO_x and NO₃), *Atmos. Chem. Phys.*, 10, 11261-11276, doi:10.5194/acp-10-11261-2010, 2010.
- 525 Reemtsma, T., These, A., Venkatachari, P., Xia, X., Hopke, P. K., Springer, A., and Linscheid, M.: Identification of fulvic acids and sulfated and nitrated analogues in atmospheric aerosol by electrospray ionization Fourier transform ion cyclotron resonance mass spectrometry, *Anal. Chem.*, 78, 8299-8304, 2006.



- 530 Rissanen, M. P., Kurtén, T., Sipilä, M., Thornton, J. A., Kangasluoma, J., Sarnela, N., Junninen, H., Jørgensen, S., Schallhart, S., Kajos, M. K., Taipale, R., Springer, M., Mentel, T. F., Ruuskanen, T., Petäjä, T., Worsnop, D. R., Kjaergaard, H. G., and Ehn, M. J.: The formation of highly oxidized multifunctional products in the ozonolysis of cyclohexene, *Am. Chem. Soc.*, 136, 15596-15606, 2014.
- 535 Robinson, A. L., Donahue, N. M., Shrivastava, M. K., Weitkamp, E. A., Sage, A. M., Grieshop, A. P., Lane, T. E., Pierce, J. R., and Pandis, S. N.: Rethinking Organic Aerosols: Semivolatile Emissions and Photochemical Aging, *Science*, 315, 1259-1262, 2007.
- 540 Sandrini, S., Pinxteren D. V., Giulianelli, L., Herrmann, H., Poulain, L., Facchini, M. C., Gilardoni, S., Rinaldi, M., Paglione, M., Turpin, B. J., Pollini, F., Bucci, S., Zanca, N., and Decesari, S.: Size-resolved aerosol composition at an urban and a rural site in the Po Valley in summertime: implications for secondary aerosol formation, *Atmos. Chem. Phys.*, 16, 10879–10897, 2016.
- Slowik, J. G., Wong, J. P. S., and Abbatt, J. P. D.: Real-time, controlled OH-initiated oxidation of biogenic secondary organic aerosol, *Atmos. Chem. Phys.*, 12, 9775–9790, doi:10.5194/acp-12-9775-2012, 2012.
- 545 Szmigielski, R., Surratt, J. D., Gomez-Gonzalez, Y., Van der Veken, P., Kourtchev, I., Vermeylen, R., Blockhuys, F., Jaoui, M., Kleindienst, T. E., Lewandowski, M., Offenberg, J. H., Edney, E. O., Seinfeld, J. H., Maenhaut, W., and Claeys, M.: 3-Methyl-1,2,3-Butanetricarboxylic Acid: An Atmospheric Tracer for Terpene Secondary Organic Aerosol, *Geophys. Res. Lett.*, 34, L24811, doi:10.1029/2007gl031338, 2007.
- 550 Tkacik, D. S., Lambe, A. T., Jathar, S., Li, X., Presto, A. A., Zhao, Y., Blake, D. R., Meinardi, S., Jayne, J. T., Croteau, P. L., and Robinson, A. L.: Secondary organic aerosol formation from in-use motor vehicle emissions using a Potential Aerosol Mass reactor, *Environ. Sci. Technol.*, 48, 11235–11242, 2014.
- 555 Tsigaridis, K., Daskalakis, N., Kanakidou, M., Adams, P. J., Artaxo, P., Bahadur, R., Balkanski, Y., Bauer, S. E., Bellouin, N., Benedetti, A., Bergman, T., Berntsen, T. K., Beukes, J. P., Bian, H., Carslaw, K. S., Chin, M., Curci, G., Diehl, T., Easter, R. C., Ghan, S. J., Gong, S. L., Hodzic, A., Hoyle, C. R., Iversen, T., Jathar, S., Jimenez, J. L., Kaiser, J. W., Kirkevåg, A., Koch, D., Kokkola, H., Lee, Y. H., Lin, G., Liu, X., Luo, G., Ma, X., Mann, G. W., Mihalopoulos, N., Morcrette, J.-J., Müller, J.-F., Myhre, G., Myriokefalitakis, S., Ng, N. L., O'Donnell, D., Penner, J. E., Pozzoli, L., Pringle, K. J., Russell, L. M., Schulz, M., Sciare, J., Seland, Ø., Shindell, D. T., Sillman, S., Skeie, R. B., Spracklen, D., Stavrou, T., Steenrod, S. D., Takemura, T., Tiitta, P., Tilmes, S., Tost, H., van Noije, T., van Zyl, P. G., von Salzen, K., Yu, F., Wang, Z., Wang, Z., Zaveri, R. A., Zhang, H., Zhang, K., Zhang, Q., and Zhang, X.: The AeroCom evaluation and intercomparison of organic aerosol in global models, *Atmos. Chem. Phys.*, 14, 10845–10895, doi:10.5194/acp-14-10845-2014, 2014.
- 560 F. Yasmeen, R. Vermeylen, R. Szmigielski, Y. Iinuma, O. Böge, H. Herrmann, W. Maenhaut, and M. Claeys, Terpenylic acid and related compounds: precursors for dimers in secondary organic aerosol from the ozonolysis of α - and β -pinene, *Atmos. Chem. Phys.*, 10, 9383–9392, 2010.
- 565 Yu, G., Bayer, A. R., Galloway, M. M., Korshavn, K. J., Fry, C. G., and Keutsch, F. N.: Glyoxal in Aqueous Ammonium Sulfate Solutions: Products, Kinetics and Hydration Effects, *Environ. Sci. Technol.*, 45, 6336–6342, 2011.
- 570 Zhang, Q., J. L. Jimenez, M. R. Canagaratna, J. D. Allan, H. Coe, I. Ulbrich, M. R. Alfarra, A. Takami, A. M. Middlebrook, Y. L. Sun, K. Dzepina, E. Dunlea, K. Docherty, P. F. DeCarlo, D. Salcedo, T. Onasch, J. T. Jayne, T. Miyoshi, A. Shimonono, S. Hatakeyama, N. Takegawa, Y. Kondo, J. Schneider, F. Drewnick, S. Borrmann, S. Weimer, K. Demerjian, P. Williams, K. Bower, R. Bahreini, L. Cottrell, R. J. Griffin, J. Rautiainen, J. Y. Sun, Y. M. Zhang, and D. R. Worsnop: Ubiquity and dominance of oxygenated species in organic aerosols in anthropogenically-influenced Northern Hemisphere midlatitudes, *Geophys. Res. Lett.*, 34, L13801, doi:10.1029/2007GL029979, 2007.
- 575 Zhang, X., Cappa, C. D., Jathar, S. H., McVay, R. C., Ensberg, J. J., Kleeman, M. J., and Seinfeld, J. H.: Influence of vapor wall loss in laboratory chambers on yields of secondary organic aerosol, *P. N. A. S.*, 111, 5802–5807, doi:10.1073/pnas.1404727111, 2014.
- 580

Tables



Sample	Oxidation level	PAM Lamp voltage (V)	OH Exposure (molec/cm ³ *s)	Collection Time (h)	average AMS OM concentration (RIE*CE=1.4) (µg/m ³)	OM on filter based on AMS (RIE*CE=1.4) (µg)
<i>α-pinene</i>						
Pin#1	High (f44 = 0.24)	110	2.10x10 ¹²	18.5	46.1	384
Pin#2	Med. (f44 = 0.11)	75	1.10x10 ¹²	3.2	151.6	221
Pin#3	Med.	75	1.10x10 ¹²	20.2	14.1	129
Pin#5	Low	30	2.00x10 ¹¹	20.5	7.5	69
Pin#6	Low (f44 = 0.05)	30	2.00x10 ¹¹	7.1	50.3	161
<i>Naphthalene</i>						
Nap#1	High (f44 = 0.30)	110	2.10x10 ¹²	19.7	31.3	277
Nap#2	Med. (f44 = 0.19)	75	1.10x10 ¹²	7	55.3	174
Nap#3	Low (f44=0.084)	30	2.00x10 ¹¹	6.6	16.9	50
Nap#4	Med (f44 = 0.20)	75	1.10x10 ¹²	6.6	42.9	127
Nap#5	Low (f44 = 0.074)	30	2.00x10 ¹¹	15.8	34.7	246
<i>blanks</i>						
Blk#1		110	2.10x10 ¹²	6.1	0	0
Blk#2		110	2.10x10 ¹²	23.2	0	0
Blk#3		110	2.10x10 ¹²	6.1	0	0

585

Table 1. PAM experimental conditions for naphthalene and α -pinene SOA ageing studies.

590

Sample	Oxidation level	PAM Lamp voltage (V)	OH Exposure (molec/cm ³ *s)	Collection Time (h)	average AMS OM concentration (RIE*CE=1.4) (µg/m ³)	OM on filter based on AMS (RIE*CE=1.4) (µg)
<i>Isoprene</i>						
Iso#1	Med (f44 = 0.046)	60	7.8*10 ¹¹	3.7	409	651



Iso#2	med	60	$7.8 \cdot 10^{11}$	2.8	575	700
Iso#3	Med (f44 = 0.039)	60	$7.8 \cdot 10^{11}$	2.2	678	656
Iso#4	Med (f44 = 0.037)	60	$7.8 \cdot 10^{11}$	2.9	551	684
Iso#5	Med (f44 = 0.040)	60	$7.8 \cdot 10^{11}$	16.1	685	4959
Iso#6	Med (f44 = 0.066)	60	$7.8 \cdot 10^{11}$	3.9	767	1280
Iso#7	Med (f44 = 0.059)	60	$7.8 \cdot 10^{11}$	3.8	593	986
<i>Blanks</i>						
Blk#1		60	$7.8 \cdot 10^{11}$	15.9	0	0
Blk#2		60	$7.8 \cdot 10^{11}$	3.1	0	0
Blk#3		60	$7.8 \cdot 10^{11}$	16.3	0	0

Table 2 PAM experimental conditions for isoprene SOA ageing studies.

595

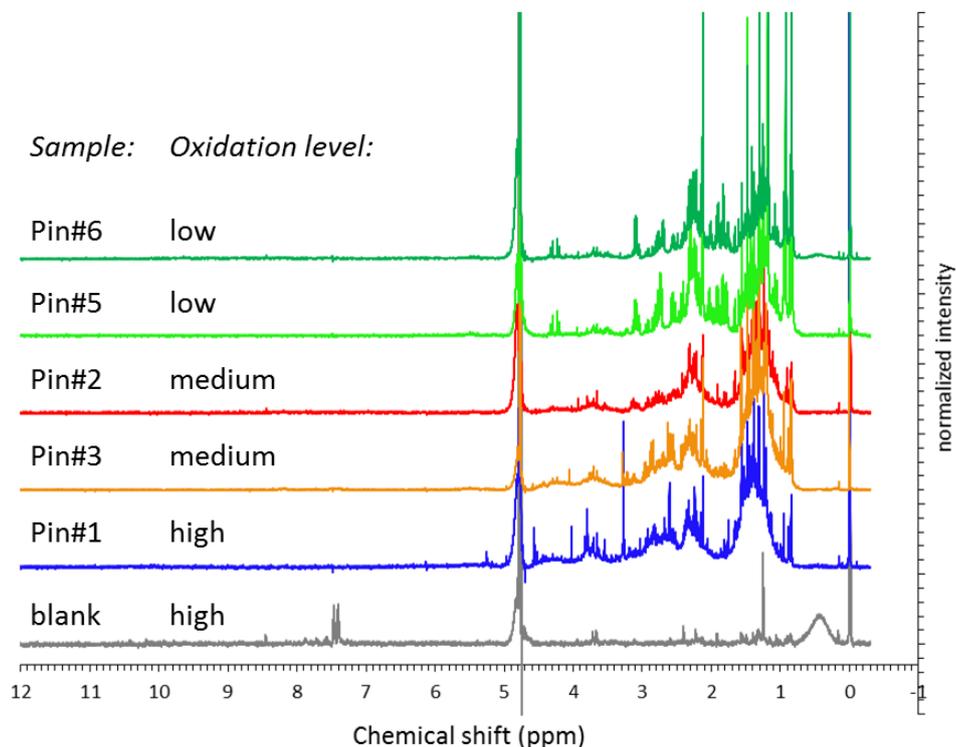
600

605

610

615

Figures



620 **Figure 1.** ^1H -NMR spectra of α -pinene SOA as a function of increasing photochemical age in the Potential Aerosol Mass (PAM) oxidation
flow reactor. The sharp singlet at zero ppm represents the internal standard (Tsp-d4), while the broad peak at 4.8 ppm is the – partly
instrumentally suppressed – HDO peak. The sharp singlets between 0.9 and 2.2 ppm in the fresh SOA samples (Pin#5 and Pin#6) are
genuine bands of the samples and were identified as methyl groups of pinic and pinonic acid.

625

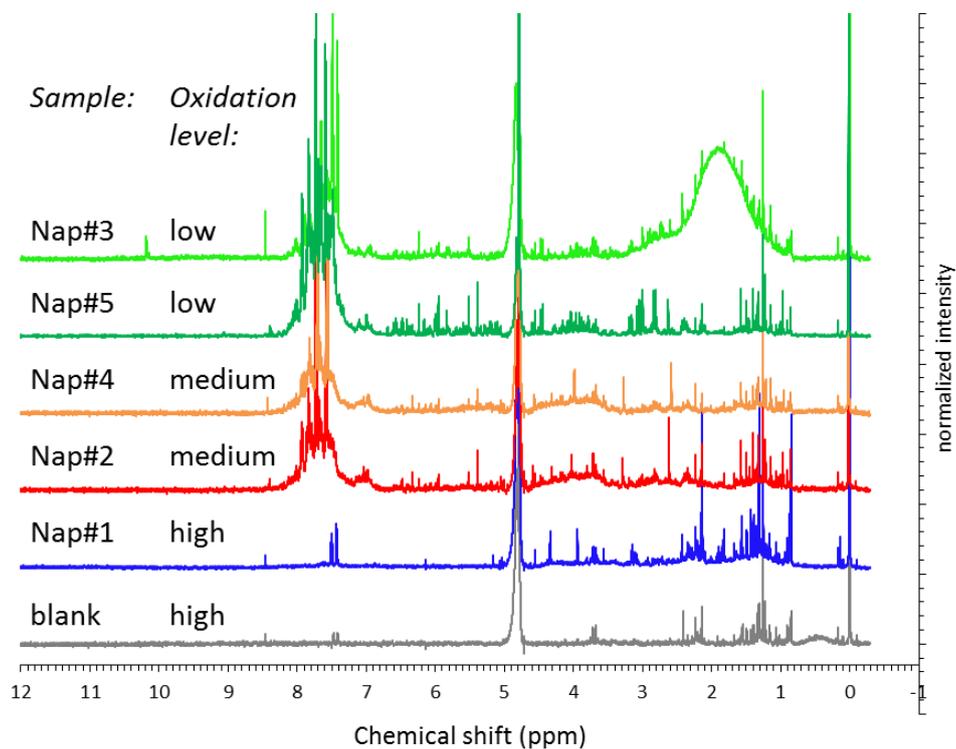


Figure 2. $^1\text{H-NMR}$ spectra of naphthalene SOA as a function of increasing photochemical age in the PAM reactor. The sharp singlet at zero ppm represents the internal standard (Tsp-d4), while the broad peak at 4.8 ppm is the – partly instrumentally suppressed – HDO peak.

630

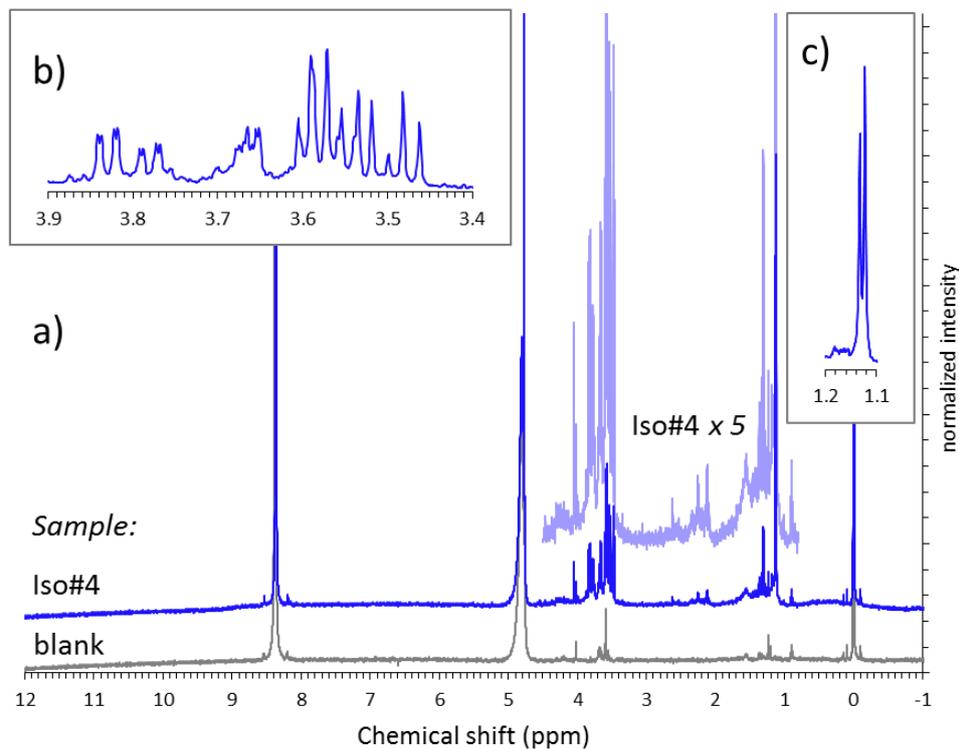


Figure 3. Panel a: $^1\text{H-NMR}$ spectrum of isoprene SOA generated in the PAM reactor at an OH exposure of $8 \cdot 10^{11}$ molec cm^{-3} sec. The bottom trace shows the same spectrum with enlarged the broad background bands. Panel b and c show the methyltetrols resonances between 3.4 and 3.9 ppm and 1.12-1.13 ppm, respectively

635

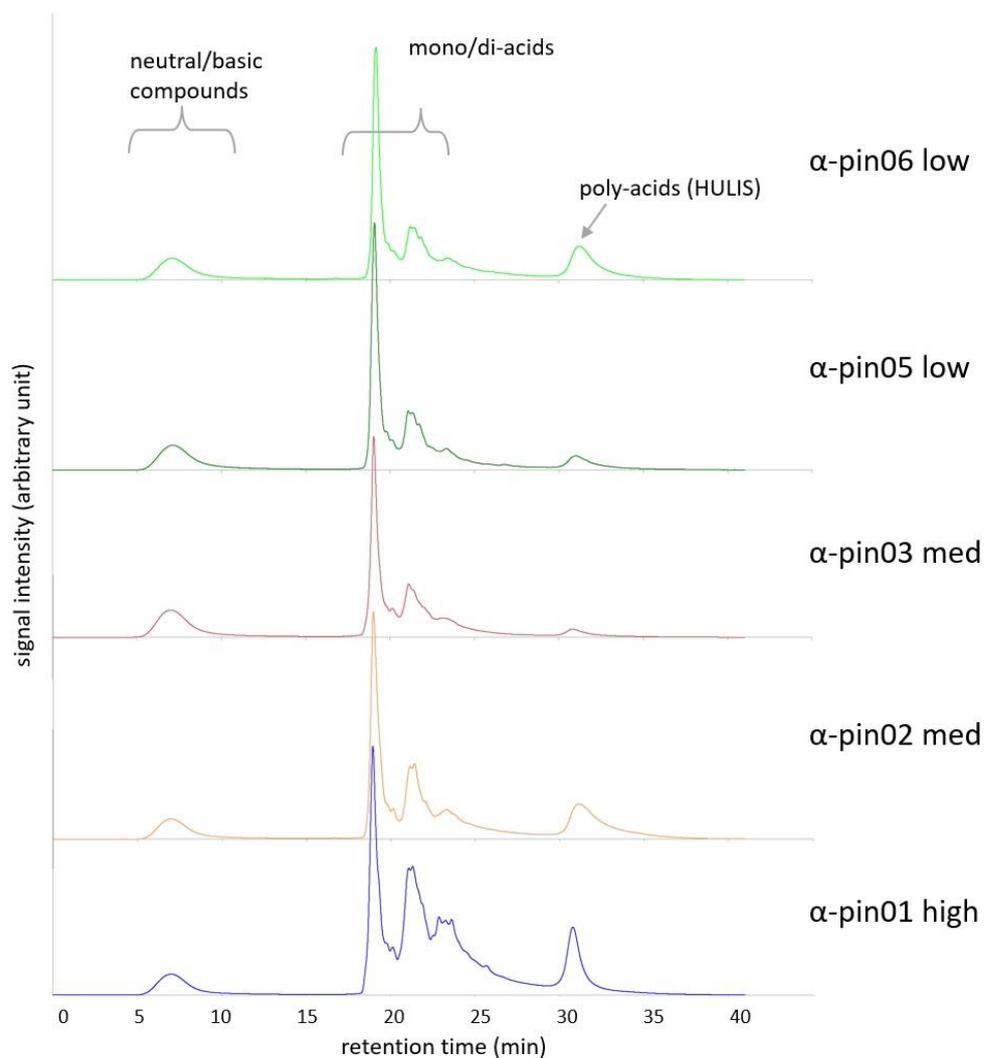
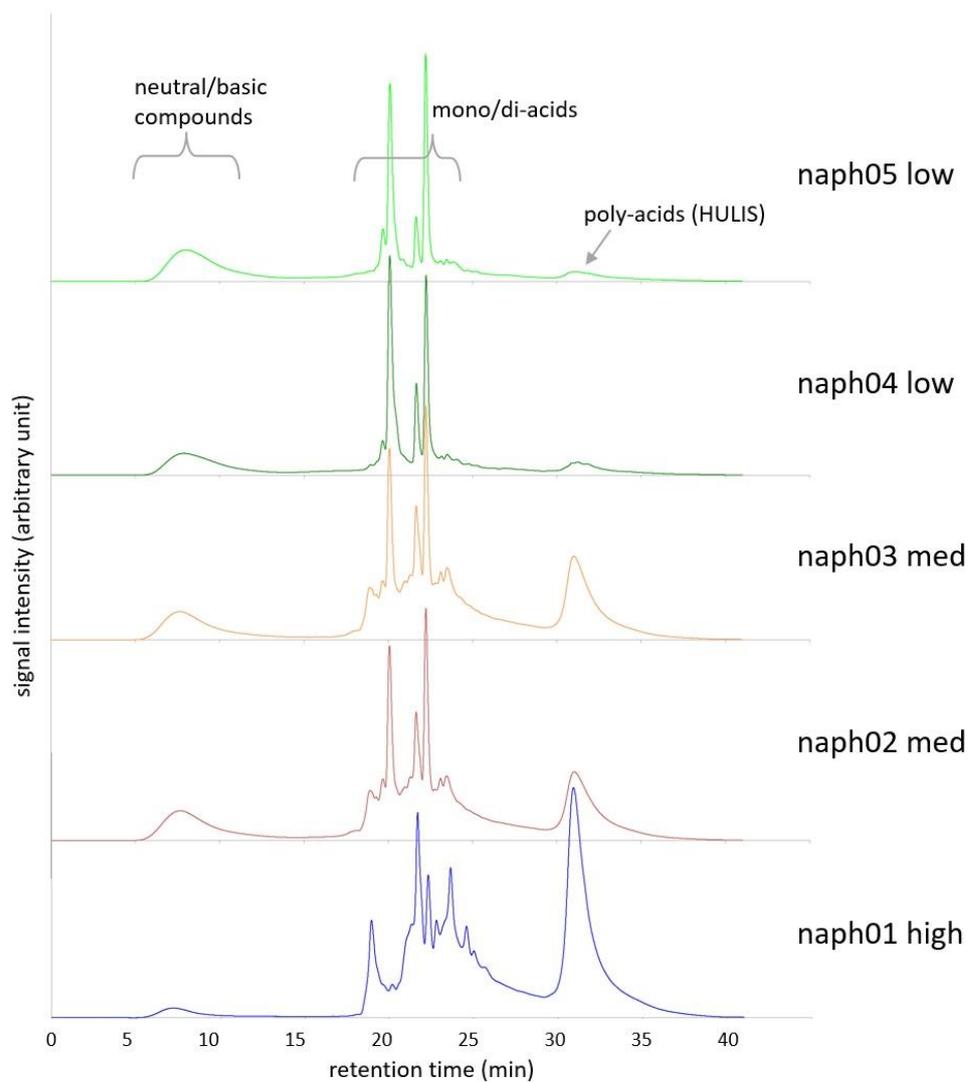
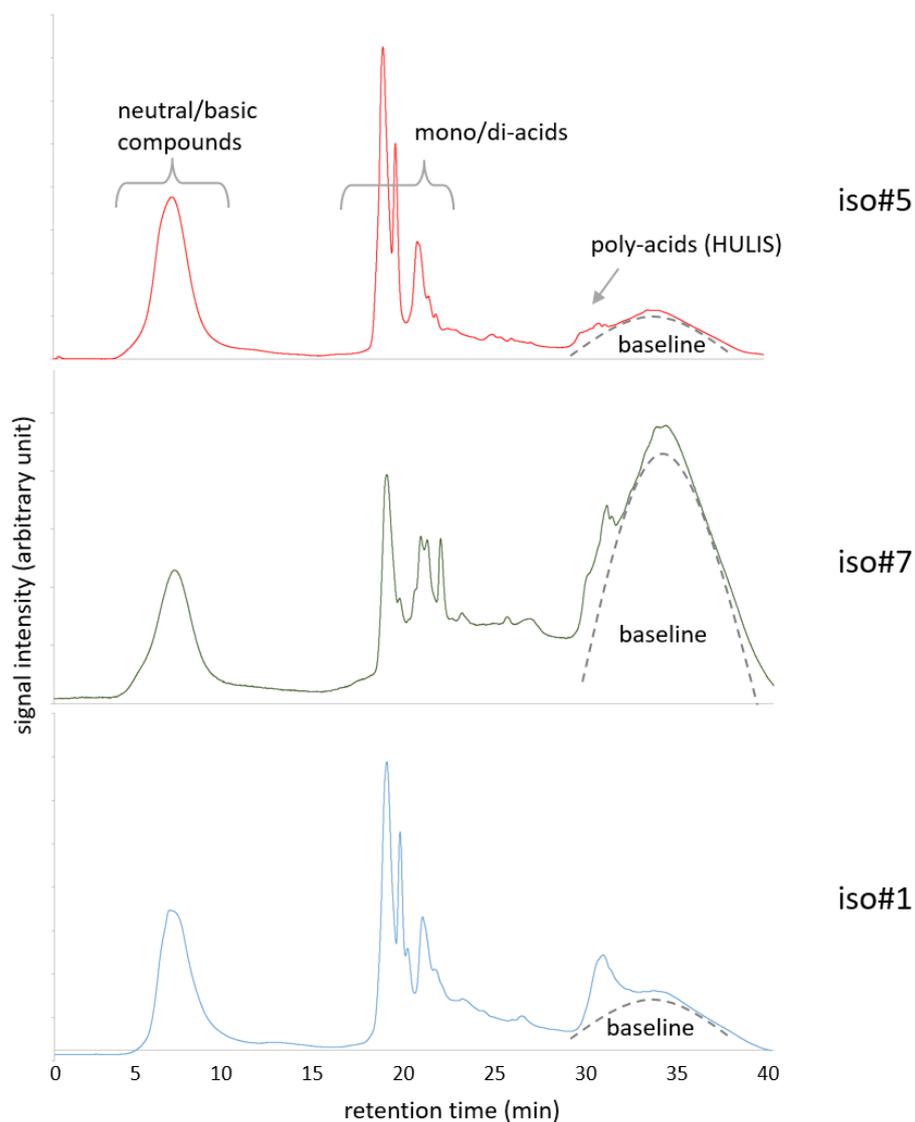


Figure 4. HPLC chromatograms of α -pinene SOA water extracts. Chromatographic features are grouped into neutral, mono-/di-carboxylic acid, and polycarboxylic acid classes based on their affinity for the column phase. Sample identifications are provided in Table 1.

640



645 **Figure 5.** HPLC chromatograms of naphthalene SOA water extracts. Chromatographic features are grouped into neutral, mono-/di-carboxylic acid, and polycarboxylic acid classes based on their affinity for the column phase. Sample identifications are provided in Table 1.



650

Figure 6. HPLC chromatograms of isoprene SOA water extracts. Chromatographic features are grouped into neutral, mono-/di-carboxylic acid, and polycarboxylic acid classes based on their affinity for the column phase. Sample identifications are provided in Table 2.

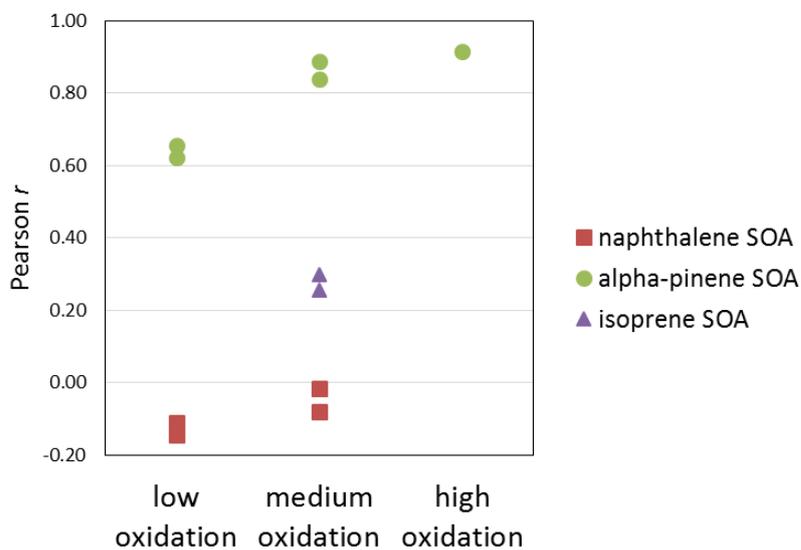
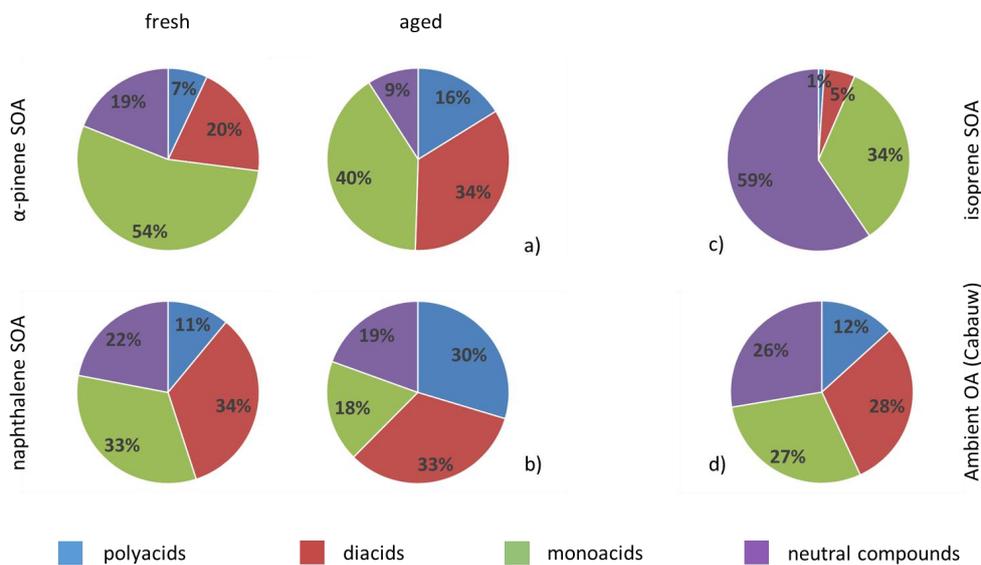
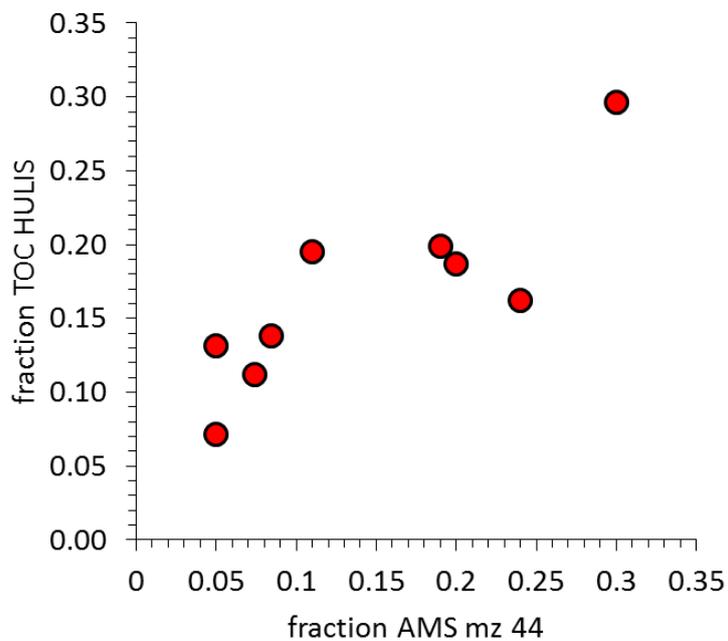


Figure 7. Pearson correlation coefficient between H-NMR spectra of PAM-generated SOA and ambient PEGASOS WSOC.



655

Figure 8. Distribution of HPLC fractions (total recovered TOC content = 100%) for α -pinene SOA (panel a), naphthalene (panel b), isoprene (panel c), and for ambient OA sampled in Cabauw, Netherlands (panel d).



660

Figure 9. Correlation plot between the AMS f44 of SOA and the HULIS fraction of HPLC-eluted WSOC.